Development and Implementation of Software Elements using State-of-the-Art Computational Methodology to Advance Modeling Heterogeneities and Mixing in Earth's Mantle Ying He^[1,3], Elbridge Gerry Puckett^[1,3], and Magali I. Billen^[2,3] UNIVERSITY OF CALIFORNIA [1] Department of Mathematics [2] Department of Earth and Planetary Sciences [3] Computational Infrastructure for Geodynamics University of California, Davis, CA 95616 USA

The Geophysical Motivation **Numerical Results**

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INFRASTRUCTURE

At the elevated pressures and temperatures of Earth's deep interior, mantle rock responds to stress by slow, creeping, solid-state flow. The resulting convection in the Earth's mantle is the driving mechanism behind subducting slabs that occur at continental boundaries and are the cause of earthquakes and volcanism. There is a very large viscosity gradient near the slab tip and, since the viscosity depends exponentially on the temperature, numerical errors in the temperature such as *overshoot* severely degrade the accuracy temperature of the model. We have developed a high-order accurate, bound preserving method for solving the temperature equation in computational subduction models.

> We studied isoviscous test cases involving a rising square with Péclet numbers from 1e3 to 1e6.

> We are now using this algorithm in ASPECT to model subducting slabs, where there is a sharp viscosity contrast near the slab tip. The sharp viscosity contrast produces thin thermal boundary layers that "act" like isoviscous

Conclusions and Future work

We approximate solutions of the incompressible Stokes equations with an advectiondiffusion equation for the temperature and a Boussinesq approximation for the density.

$$
-\nabla \cdot [2\eta \epsilon(\mathbf{u})] + \nabla p = \rho \mathbf{g},
$$

$$
\nabla \cdot \mathbf{u} = 0,
$$

$$
\frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = \nabla \cdot (\kappa \nabla T),
$$

$$
\rho = \rho (1 - \rho (T - T))
$$

 $\rho = \rho_0 (1 - \alpha (1 - I_0)).$ The temperature equation (3) is convection-dominated, if the Péclet number $P_e = \frac{L\|\mathbf{u}\|_{\infty}}{E} >> 1$. Note, if $\kappa = 0$, there is no diffusion and $Pe = +\infty$.

The following two advection methods are implemented in ASPECT

Discontinuous Galerkin Method (DG)

Finite Element Method (FEM)

Model Problem

The original advection-diffusion algorithm is an *FEM algorithm with entropy viscosity stabilization*

We have developed a DG advection algorithm with a monotone flux and limiting

• Our numerical results have demonstrated that, compared to FEM with entropy viscosity, DG with limiting is more suitable for convection dominated flows; i.e., higher

-
- Péclet numbers.
- flows with large Péclet numbers.

[1] Y. He, E.G. Puckett, and B.I. Magali, 2017, A Discontinuous Galerkin Method with a Bound Preserving Limiter for the Advection of non-Diffusive Fields in Solid Earth Geodynamic, Physics of the Earth and Planetary Interiors [2] Kronbichler, M., Heister, T., Bangerth, W., 2012. High accuracy mantle convection simulation through modern numerical methods. Geophysical Journal International [3] J.L. Guermond, and R. Pasquetti and . Popov. 2011 , Entropy viscosity method for nonlinear conservation laws. Journal of Computational Physics [4] Zhang, X., Shu, C.-W., 2010. On maximum-principle-satisfying high order schemes for scalar conservation laws. Journal of Computational Physics 229 (9)

Crustal layers for slab dynamics, *Arredondo & Billen, J. Geodynamics (submitted)*

ASPECT (Advanced Solver for Problems in Earth's ConvecTion) is an open source C++ code to model problems in the Earth's mantle. https://aspect.dealii.org/

 (2) (3)

 (4)